1	Spatiotemporal quantification of key environmental changes in Stok and
2	Kang Yatze regions of Ladakh Himalaya, India
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30 Abstract:

Tourism fuelled economic transformation of Ladakh region is bringing a change in demography and livelihood preferences leading to increased exploitation of scarce natural resources. In addition to the nerve centre of tourism-related activities (Leh city), rural hinterlands of Ladakh are also witnessing a stark transformation in land and resource use patterns. In this study we quantified and assessed the changes in key environmental parameters of rural Ladakh between 1969 and 2020. We observed a decrease in glacier area, glacier volume and vegetation cover by ~19%, 13% and 14%, whereas built-up and lake area increased by ~238% and 122%, respectively. Simultaneously, local population of the region doubled while tourism footprint increased by ~620 folds. The study highlights the environmental changes happening at a higher rate in rural regions around Leh city. Therefore, need of the hour is to recognise the disruptive changes being brought in the traditional way of life and to devise strategies to reorient the development paradigm towards sustainable resource use and management.

45 Keywords: Ladakh, tourism, environment, climate change, glacier, Himalaya.

59 **1. Introduction**

Mountainous regions in Himalaya face various sustainability challenges due to climate, 60 61 geopolitical, economic, and societal transformations, which affect environment, water resources and demography of the land (Bhardwaj et al. 2019). Some of the major factors for 62 such transformations are the climate change (Sam et al. 2019), booming tourism industry, 63 international conflicts, rapid urbanisation and migration. These factors will have both positive 64 and negative impact on the environment and society. Positive impacts comprise of warmer 65 climate more conducive to habitation, improvement in living standards and economy, 66 agricultural growth, and improved connectivity, among others. However, the negative impacts 67 are often long-lasting, irreversible, and more damaging, and consist of excessive stress on the 68 environment, unplanned development, exploitation of natural resources, and pollution, among 69 others. Policy design and implementation is challenging in mountains, because usually these 70 regions have been of marginal concern due to lesser population, and therefore often ignored in 71 development priorities (Messerli and Ives 1997; Geneletti and Dawa 2009). Environmental 72 fragility and tourism seasonality of mountain regions further add to the issues hindering 73 systematic and integrated developments (Buckley et al. 1999; Arrowsmith and Inbakaran 2002; 74 75 Geneletti and Dawa 2009). Moreover, mountain areas often mark international borders for many countries and thus, hold political sensitivity in terms of major policy changes (Nepal and 76 77 Chipeniuk 2005). With increasing globalisation and political stability in majority of the world, mountain tourism has developed with an accelerated pace in the past several decades (Price 78 1992; Moss and Godde 1999). However, such tourism has led to unprecedented growing 79 environmental concerns in developed (Draper, 2000) and developing countries alike (Stevens, 80 81 2003, Buntaine et al., 2006).

One of such mountainous regions is Ladakh which used to be a restricted land for the outside world due to its geo-strategic sensitivity (Bray, 1991; Warikoo et al., 2020) and tribal

population (Smith, 2012). However, the government of India opened the region for tourism in 84 1974 (Bray, 1991), which marked the beginning of a new era for Ladakh. As per the data 85 available from the Leh tourist office, 8,608 Indian and 16,256 foreign tourists visited Ladakh 86 in 1988 (Bray, 1991). Thus, tourism emerged as the prime revenue generator for the region and 87 opened new income sources for guesthouses and hotels, local shops, and guided tours and 88 travels (Bray, 1991; Nüsser, et al., 2019a,b; Müller et al., 2020). However, the initial economic 89 90 benefits of tourism had been concentrated in the Indus valley, disproportionately benefitting non-Ladakhis (Michaud 1991). With improvements in connectivity 91 mainly and 92 communication, this scenario has evolved in the past two decades as the number of visitors increased many folds over last few decades from a few hundred in 1974 to more than 300,000 93 visitors in 2019 from every corner of the world (Muller et al., 2020). Agriculture used to be the 94 95 primary sector for the local economy but today it is replaced by tertiary sector (especially tourism) (Dame et al., 2019; Muller et al., 2020). Leh, which is the capital city of Ladakh, has 96 now become one of the most visited cities in high-Himalayan Mountains. The recent decision 97 to develop Ladakh as a Union Territory of India has cleared the path for more rapid urbanisation 98 and development of the region in coming years. 99

100 While rapid urbanisation of Leh city has recently been investigated from a remote sensing perspective (Dame et al., 2019, Müller et al., 2020), the effects are still unknown for the rural 101 areas around Stok and Kang Yatze range. As mentioned above, such regions along the Indus 102 valley have been the target of this tourism-led changes in Ladakh (Michaud 1991; Nüsser et 103 al., 2012). Thus, for these regions, it is particularly important to quantify the changes in several 104 of the key environmental parameters such as glaciers, vegetation, water, and built-up, between 105 106 1969 and 2020, to understand the long-term impacts of tourism and demographic changes on high-mountain catchments. In this study we are trying to understand and document the major 107 land cover changes happened in the parts of rural Ladakh for the past 50 years (1969-2020). 108

109 The main objectives of this research are: (1) to perform temporal mapping of glaciers, 110 vegetation, water, and build-up in Stok and Kang Yatze regions of Ladakh Himalaya; (2) to 111 quantify spatiotemporal changes in the mapped parameters; and (3) to use climate, 112 demography, and other ancillary datasets for further assessing the quantified changes.

113 **2.** Study area

In this study, two different adjacent regions (Stok and Kang Yatze region) with various 114 glaciers, build-up, vegetation, and population settings were selected from Ladakh, India (Figure 115 1) to understand the changes occurred due to climate change, tourism, and urbanisation during 116 past five decades. The geographic extent of the study area lies within latitude and longitude of 117 33.3° to 34.3° N and 77.1° to 78.1° E, respectively. It covers a vast area around Stok and Kang 118 Yatze regions extending from 3000 to 6400m a.s.l in the northern Zanskar range, Ladakh, India. 119 120 The region of interest consists of 15 villages of various sizes with population ranging from 128 - 5,565 individuals (Census 2011). Most of these villages are entirely dependent on the glacier 121 and snow meltwater for irrigation and domestic use. Stok and Kang Yatze regions house 30 122 and 106 glaciers of various sizes ranging from 0.01 to 5 km², with larger glaciers (>3 km²) 123 entirely concentrated in Kang Yatze region. In 2019 the total glacierised area was around 89 124 km^2 , equivalent to 3.2% of the study area. 125

Ladakh, the newly formed Union Territory of India, has a unique climate with extremely arid and cold environment. As per the longest in-situ data available from the region headquarter Leh (3500m a.s.l), the long-term yearly mean precipitation sum is <100 mm whereas the mean annual temperature is ~5.6 °C (Nüsser et al., 2012; Soheb et al., 2020). The region is also a hub for wildlife and climbing enthusiasts for its diverse flora fauna and the two famous peaks of Stok (6140m a.s.l.) and Kang Yatze (6400m a.s.l.) summit. Apart from agriculture and service sector, tourism has developed as a major driver of the local economy of the region in the current times (Dame et al., 2019; Müller et al., 2020). Figure 1 presents the study area and the keyenvironmental parameters within the region.

135 **3. Data and methods**

In this study, we used freely available multiple datasets from various sources to understand the transformations in key environmental parameters such as vegetation, build-up, glacierised areas, water bodies, and demography (Table 1). The information on multisource datasets used within this study are provided in Table 1. Below we provide descriptions of used datasets and methods for each of the objectives.

141 *3.1.Glacier mapping*

For glacier mapping, Orthorectified level-1T Landsat-Thematic Mapper (TM) and Operational 142 Land Imager (OLI) images (www.earthexplorer.usgs.gov) with minimum (<30%) cloud cover 143 144 from ablation period (August-September) were used to map the glacierised region of different time periods (1991, 1998, 2010 and 2019). Several robust and established methods to map 145 glaciers are available in literatures e.g. Band ratio approach (Paul et al., 2002, 2015; 146 Racoviteanu et al., 2009; Bhardwaj et al., 2014; Bhardwaj et al., 2015; Schmidt & Nüsser, 147 2017; Smith et al., 2015; Winsvold et al., 2014, 2016), supervised classification methods 148 (Muhammad et al., 2013; Khan et al., 2015; Nagai et al., 2016; Tian et al., 2017; Muhammad 149 2019), and manual delineation method among others. 150

In this study, glacier mapping was carried out by closely following the Global Land Ice Measurements from Space guidelines (Raup et al., 2010). We used the ratio approach between Near Infrared and Shortwave Infrared bands with a threshold of 2.0 (NIR/SWIR > 2 = ice/snow). A median filter of kernel size 3 x 3 was applied to remove the isolated and small pixels outside the glacier area. The NIR and SWIR band ratio approach is good at distinguishing glacier pixels from water features with similar spectral reflectance values (Racoviteanu et al. 2009; Zhang et al. 2019). Band ratio technique doesn't perform well in
shadowed regions therefore additional manual corrections was required (Muhammad and Tian
2016). This method is robust and time-efficient for mapping clean glacier (Paul and Kaab,
2005; Andreassen et al., 2008) and fortunately the study area only consists of clean glaciers.

Multiple high resolution PlanetScope images (www.planet.com) of 29 August 2018 and Google Earth were also used to improve the mapping for the current extent of glaciers. The extent of glaciers in the accumulation zone were kept unchanged throughout the mapping periods because no significant changes is expected in these regions, and also this approach helps in minimising the error in mapping a glacier (Bolch et al., 2010; Bhambri et al., 2013; Garg et al., 2019).

Six Corona KH-4A images (www.earthexplorer.usgs.gov) from 30 July 1969, with minimum 167 168 snow cover, were used to extract the historical extent of glaciers of the study area. Owing to the complex geometry of corona imagery, 12 subsets were made for the study area. All the 169 subsets were separately co-registered following the steps performed by (Bhambri et al., 2011). 170 A projective transformation using ground control points (GCPs) and the Advanced Spaceborne 171 Thermal Emission and Reflection Radiometer Digital Elevation Model (ASTER DEM) was 172 performed followed by a spline adjustment in ArcGIS 10.5. Stable boulders, stupas, historical 173 sites, Monasteries, road junctions and bridges among others were used to collect GCPs. The 174 175 primary focus of the Corona image rectification was to georectify the glacierised and build-up areas. Digitization of 1969 glacier outlines were manually carried out on the corrected Corona 176 image subsets. 177

178 **3.2.** *Glacier thickness and volume estimation*

ASTER DEM from two periods at a temporal gap of atleast 10 years, with no cloud cover overglaciers of the study area, from the ablation time was required to estimate the volume of these

ice masses of the study area and see the changes over time. DEM from 2004 and 2020 satisfied 181 the conditions and were chosen for further analysis. The ASTER DEM products (AST14DMO) 182 are created using only orbital ancillary data and without GCPs, therefore small outliers (peaks 183 and sinks) are likely to be present in DEMs. These outliers were removed using the method 184 followed by Shukla and Garg, 2019. Relative horizontal and vertical biases are often present 185 in DEMs derived from different sensors (Nuth and Kaab, 2011). In the present study, we 186 187 observed a slight shift in the DEMs from the same sensor for the two periods (2004 and 2020). Therefore, the three-dimensional coregistration of the DEMs was required to be performed on 188 189 stable areas (least error prone regions, e.g., flat valleys, grass lands and farms among others). We followed the coregistration method developed by (Nuth and Kaab, 2011) to minimise the 190 horizontal and vertical biases. The glacierised areas were first removed from the DEMs 191 followed by the areas with a slope between 4-45 degrees, cloud cover, large outliers (elevation 192 difference $>\pm 100$ m) and pixel elevation below 4000m a.s.l. (as no glacier is present in this 193 region below this elevation). The horizontal shift between the reference and slave DEM was 194 corrected significantly. The X and Y shift was reduced from 26.5 and 23.81 m to 4.13 and 3.45 195 m, respectively. However, the vertical biases available in the DEMs were corrected using all 196 (> 20,000 pixels) the reliable elevations differences over the stable terrains (Nuth and Kaab, 197 2011). Furthermore, we have used these temporal bias-corrected DEMs as input in GlabTop 198 and GlabTop2 models to estimate glacier volumes rather than change in surface volume (or 199 200 DEM differencing). The models have their own uncertainties, as discussed in the later section, which are well-above any jitter-induced errors and did not make any notable difference in the 201 GlabTop model results. 202

The bias corrected DEMs were further used to estimate the volume of glaciers of the study area. We estimated the volume using three different methods, i.e., Glacier Bed Topography model (GlabTop2), shear-stress/slope dependent (hereafter slope dependent method) and area-

related methods. GlabTop2 model is the upgraded version of the previous GlabTop model 206 developed by (Linsbauer et al., 2009) where the thickness is estimated at several points along 207 the glacier branch lines further interpolating to the entire glacier (Linsbauer et al., 2009; Paul 208 and Linsbauer et al., 2012; Frey et al., 2014). In the advanced GlabTop2 model, the thickness 209 is calculated for randomly selected DEMs pixel within the glacier inner cells (or glacierised 210 area). Thickness of all other glacier cells are estimated by interpolating the thickness at random 211 cells and the ice thickness at the glacier marginal cells (or along glacier margins) where ice 212 thickness is known to be zero (Frey et al., 2014; Ramsankaran et al., 2018). The calculation is 213 214 distributive, requires only DEM and glacier outline as input and uses Equation 1 (Cuffey and Paterson, 2010; Haeberli and Hoelzle, 1995). 215

216
$$hf = \frac{\tau}{f\rho g \sin(\alpha)}$$
 (1)

217 where

218 τ = average basal shear stress along the central flow line

f = shape factor (60-90; Cuffey and Paterson, 2010) f = 1000

220 $\rho = \text{density of ice } (900 \text{ kg m}^{-3})$

221 g = acceleration due to gravity (9.8 m s⁻²)

222 α = mean surface slope

223

The value of basal shear stress (in kPa) is estimated using an empirical relationship between basal shear stress and glacier elevation range (Δ H, in km) as per Equation 2 (Haeberli and Hoelzle, 1995)

227
$$\tau = \begin{pmatrix} 0.5 + 159.8 \,\Delta H - 43.5 \,(\Delta H)^2 & \Delta H \le 1.6 \,km \\ 150 & \Delta H > 1.6 \,km \end{pmatrix}$$
(2)

Slope dependent method uses the same equation (1 and 2), as Glabtop and Glabtop2 model, to 229 calculate the average thickness of a glacier. To account the semi-elliptical cross-sectional 230 geometry of a glacier, a multiplication with 0.785 is added to equation 1. This method has been 231 applied extensively by many (Hoelzle et al., 2007; Paul and Svoboda, 2009; Salzmann et al., 232 2013). However, area-related thickness estimation only uses surface area of the glacier and 233 other scaling parameters. It is the most frequently used method because of its simple and fast 234 235 implementation. Equation 3, given below, is used for area-related thickness estimation.

- $H = cA^{\gamma}$ (3) 236
- Where 237

H = mean thickness of glacier 238

c and γ = scaling parameters given by Chen and Ohmura (1990) who used 63 glaciers and Bahr 239 et al. (1997) who used theoretical analysis to determine these parameters. 240

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The obtained thickness from slope-dependent and area-scaling method was then used to 242 calculate the volume of glaciers with the help of glacierised areas of the respective years. The 243 pixel-based thickness obtained from the GlabTop2 model was first applied to calculate the 244 volume of each pixel and further summing up all the pixel to get the volume of individual 245 glacier volume. 246

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3.3. build-up, vegetation, and water body mapping

Both Corona and PlanetScope images were used to map the built-up, vegetation and water 249 bodies of 1969 and 2018, respectively. Owing to lesser heterogeneity of the landscape, a 250 supervised classification technique on 1969 and 2018 images was used to distinguish between 251 vegetation and non-vegetation land in ArcGIS 10.4, with additional manual correction where 252

needed. The classification process included 15 training samples containing >30,000 pixels 253 combined to improve the classifier's overall performance. Built-up area of 2018 were manually 254 digitised on screen using high resolution PlanetScope image. The 2018 results were overlaid 255 on the 1969 corona image as a reference to digitise the built-up area of 1969. The built-up areas 256 of 2018 which were not present or cannot be identified on 1969 images were removed from 257 1969 built-up areas. The built-up areas which no longer exist at present but were present in 258 259 1969 also mapped and added. Water-bodies were mapped manually on screen on both Corona and PlanetScope images. 260

261 *3.4.Climate data analysis*

In addition to the satellite imageries, long-term ERA5 monthly single level reanalysis climate 262 data have also been used in the study to understand the climatic settings of the study area. 263 ERA5 (downloaded from https://climate.copernicus.eu/) is a 5th generation European Centre 264 for Medium-Range Weather Forecasts (ECMWF) reanalysis high resolution climate data at 265 0.25 degrees. The data carries a large set of variables from 1979 to present. In this study, 266 monthly mean temperature and precipitation data was used. In-situ data from Leh, India 267 (nearest in-situ data available) was used to bias correct the ERA5 temperature and precipitation. 268 Kanda et al., 2020 evaluated the performance of seven grid-based datasets against the observed 269 data from 19 in-situ stations across the region of Karakoram, Greater and Lower western 270 Himalaya and found that ERA-Interim (precipitation), and CRU-TS and ERA-Interim 271 (temperature) are the best performing datasets in all these regions with additional bias 272 corrections. ERA5 is the improved version of ERA-Interim and can be used for further analysis 273 (Figure 4). 274

Linear scaling bias correction (Ines and Hansen 2006; Teutschbein and Seibert 2012, Shrestha
et al 2016) was adopted for ERA5 data correction with the help of Leh in-situ data. The bias

corrected Leh data was statistically analysed to understand the magnitude and significance of
trends using Mann-Kendall test and Sen's slope estimator (Sen, 1968), respectively. The
change in temperature and precipitation (Table 2) was calculated using Equation 4 following
Shukla et al 2020:

281 $Change = (\beta \times L)/M$ (4)

282 Where β is Sen's slope estimator, L is the length of the period and M is the long-term mean.

283

284 **3.5.** Uncertainty assessment

Our study involves extraction of various key environmental parameters utilizing multiple satellite data and methods. Hence, there are different sources of uncertainties related to the approaches used in the present study. Therefore, in this section we present the methodologies that were used to estimate the uncertainties.

289 *3.5.1.* Uncertainties in glacierised area and volume:

Since there is no available in-situ reference data for large set of glaciers in this region. Therefore, we estimate the error term based on a buffer based assessment for each glacier following the methods of Bolch et al., 2010; Granshaw & G. Fountain, 2006; Mölg et al., 2018; Tielidze & Wheate, 2018. The buffer width was set to half-pixel as the glaciers of the study area are debris free. The half-pixel buffer was applied to the glacier polygons and area was recalculated. The overall uncertainty was ~ $\pm 3.4\%$ with individual uncertainties of \pm 0.6, 3.9, 4.0, 4.1 and 4.2% for 1969, 1991, 1998, 2010 and 2019 images (Table 3).

For GlabTop2 and Slope dependent method the uncertainties were estimated based on error propagation method (Kumari et al., 2021; Ramsankaran et al 2018; Frey et al., 2014) by modifying the scaling parameters. Since the variation in these parameters are unknown for the

study area, $\pm 5\%$ of uncertainty was considered for τ , f (Driedger and Kennard 1986; Haeberli 300 and Hoelze, 1995; Kumari et al 2021; Ramsankaran et al 2018) and $\pm 10\%$ was considered for 301 the density of ice (ρ) following Gantayat et al. (2014). The uncertainty in the term u (sin α)/sin 302 α was assumed to be ± 0.2 considering the potential uncertainty of ~20m in ASTER DEM in 303 the Himalayan region (Bhambri et al., 2011; Toutin 2008). The uncertainty in the area-related 304 volume estimation was determined by comparing the influence of the two applied scaling 305 306 parameter sets on the results with additional modification of $\pm 5\%$ in individual glacier area (Frey et al., 2014). Table 4 present the associated uncertainties. 307

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309 3.5.2. Uncertainties in other environmental parameters:

310 The uncertainty in the build-up and lake area was estimated based on a comparison between the independently generated outlines in two sample villages (Chuchot and Stok) and all the 311 lakes by three different analyst on the same satellite images. For vegetation cover, a stratified 312 random sampling design was used to determine the sample numbers and their position on the 313 map. For the purpose we have selected ~500 points on both Corona and PlanetScope images. 314 315 The accuracy assessment was done by determining the probability of a particular cell value being similar to the actual generated classified information (i.e. vegetation cover). Error matrix 316 (also known as confusion matrix or covariance matrix) and kappa coefficient was calculated 317 for both Corona and PlanetScope. Details of the uncertainties are presented in Table 5. 318

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4. Results and discussion

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4.1.Overall change in key environmental parameters between 1969 and 2021

We analysed several key environmental parameters in the study area, i.e., glacierised area and volume, built-up area, vegetation area, lake area, and demography between 1969 and 2021. Due to limitations in temporal datasets, the time periods for the analyses are different. The glacierised region was mapped for 1969, 1991, 1998, 2010 and 2019. Glacier volume was estimated for 2004 and 2020 due to non-availability of cloud-free DEMs before 2004. Builtup, vegetation and lake areas were analysed for two periods (1969 and 2018) because publicly no high-resolution images exists in between. The population data is also not available after census 2011, therefore population of 2021 was extrapolated through a regression developed using available datasets of from 1971, 1981, 2001 and 2011 census reports.

330 *4.1.1. Overall change in glacierised area and volume*

Stok and Kang Yatze watersheds consist of 30 and 106 glaciers, larger than 0.01 km², covering 331 an area of 10.6 and 78.5 km², respectively. These glaciers retreated over time with varying 332 retreat rates during different periods. Glaciers of Stok region retreated at a rate of 0.26, 0.35, 333 0.6, and 0.57% per year during the periods 1969-91, 1991-98, 1998-2010, and 2010-19, 334 respectively. Total change observed over 50 years was 2.51 km², equivalent to 19% 335 (0.38%/year) of the total glacierised area. Whereas, Kang Yatze region observed a slightly 336 different rate of change during individual periods. However, total change (in %) in glacierised 337 area over 50-year period was similar to Stok region, i.e., 18.7 km², equivalent to 19% 338 (0.38%/year). The retreat rate in Kang Yatze was found to be 0.38, 0.28, 0.62, and 0.33% per 339 year for 1961-91, 1991-98, 1998-2010, and 2010-19, respectively. Overall, glacierised area 340 changed from 110.4 km² in 1969 to 89.21 km² in 2019 with a change of 21.2 km², equivalent 341 to 19% (0.38%/year) of the glacierised area over 50-year period. These changes are however 342 higher than the changes observed by Negi et al., 2021 and Shukla et al, 2020 in the adjacent 343 region of Shayok (Karakoram region) and Suru (Kargil region) where the retreat was reported 344 to be 7.8% (0.008%/year) and 6% (0.2%/year), respectively. Negi et al., 2021 studied 569 345 glaciers of size greater than 1 km² of Shayok region between 1991 and 1994. Whereas, Shukla 346 et al., 2020 studied glaciers larger than 0.01 km² between 1971 and 2017. However, Schmidt 347 and Nüsser, 2017 studied glaciers (> 0.03 km^2) of Stok, Kang Yatze and the region nearby 348

between 1969-2016 and found results similar to the present study. These results are interesting
considering that our study area is the most populous and tourism-prone within Ladakh region,
and the significantly higher glacier retreats clearly point towards a plausible connection
between glaciers and threatened microclimate. Figure 2, 3, 5 and 8 presents the area changes
over 50-year period while Figure 4 provides the data on climate changes in this region.

Two (GlabTop2 model and area-scaling method) of the three methods used in this study to 354 estimate the thickness of these ice masses gave similar results; however, the third method 355 (slope-dependent method) gave comparatively lower thickness estimate. Volume of glaciers in 356 Stok region using GlabTop2, area-scaling and slope-dependent method was found to be 0.29, 357 0.31 and 0.22 km³, respectively, for the year 2004; and 0.23, 0.27 and 0.17 km³ for the year 358 2020, respectively. Volume of glaciers in Kang Yatze region using GlabTop2, area-scaling and 359 slope-dependent method was found to be 3.2, 3.1, and 2.1 km³, respectively, for the year 2004 360 and 2.6, 2.8, and 1.7 km³ for the year 2020, respectively. Figure 2, 3, 5 and 8 present the average 361 values of these estimates over 16-year period. The average (of the three different estimates) 362 change found over the 16-year period was 0.05 and 0.37 km³, equivalent to 18% and 13% (1.1 363 and 0.8%/year) in the total ice reserve of Stok and Kang Yatze region, respectively. The study 364 area currently holds ~2.5 Gigaton (Gt) of water with 0.21 and 2.5 Gt of water in the glaciers of 365 Stok and Kang Yatze regions, respectively. 366

Changes in the glaciers of high-altitude Himalayan region are generally due to the rise in temperature, orientation, debris and black carbon among others. However, in the region around Leh, these changes are also associated with the lower winter precipitation (Soheb et al., 2020). Region around Leh (study area) generally receives comparatively less precipitation (Figure 4) (Soheb et al., 2020; Schmidt and Nüsser 2017), and together with the rise in temperature the impact will be comparatively higher on these glaciers. It is evident that the glacierised region around Leh is losing mass at a higher pace despite being in a higher altitude region between

Karakoram and other western Himalayan region. The higher retreat can be due to the presence 374 of smaller glaciers in high numbers in Ladakh (Schmidt and Nüsser, 2017). These glaciers are 375 more prone to rise in temperature and decline in solid precipitation due to their short response 376 time, and they are most probably vulnerable to local climatic variations. Leh showed an 377 increment of 0.04, 0.05 and 0.03 °C/year in annual, JJAS and winter temperature, respectively, 378 at p<0.05 (95% percent confidence level) through Man-Kendal test and Sen's slope estimator 379 380 (Sen, 1968). However, the trend estimated in precipitation is negative but not statistically significant to conclude any inferences from it (Table 2). 381

The decline of glaciers in such environment is of serious concern as it directly impacts 382 livelihood, ecosystem and socio-economic dimensions of the region because of their higher 383 dependency on glacier meltwater streams. The residents of the region have witnessed a decline 384 in agricultural yields, the main driver of economic development of the region, due to a decrease 385 in water resources (Barrett and Bosak 2018). The water scarcity together with the recent boom 386 in tourism activity (tourism footprint in Leh was 327,366 in 2018, a number that is more than 387 the entire population of Ladakh) has led to a shift in livelihood from agriculture to other 388 commercial activities (Müller et al. 2020), though even the latter relies heavily on water 389 390 resources. To cope with water scarcity, some people of Ladakh have developed new water management techniques, commonly known as 'ice reservoirs' or 'ice stupas', to supplement 391 392 agricultural activities (Nüsser, et al., 2019a,b).

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4.1.2. Overall changes in Vegetation, built-up areas and lakes

395 Vegetation cover of within the village vicinity of a catchment was analysed in this study. Two396 entire glacier catchments of Kang Yatze region were also excluded from this analysis because397 no settlement exist in those remote catchments. Vegetation cover of overall study area is

398 presented in Figure 5, 6, 7 and 8 along with the % change in 49 years. Four villages namely 399 Chuchot, Stakna, Changa and Hemis are also included in the Stok region because they are 400 partially fed by the meltwater coming from the Stok range. However, Hemis village is the only 401 village in this study which is not fed by any glacier. The source of water for Hemis village is 402 snowmelt and most probably springs within the catchment as the catchment does not have any 403 glacier but comprise one of the oldest Monastery (Hemis monastery), village and vegetation.

Stok region consists of several villages in which Rumbak, Stok, Matoo, Martselang and Shang 404 are completely fed by the glaciers of this region. Whereas, all the villages in Kang Yatze region 405 i.e., Upshi, Miru, Lato, Gya, Sa-Soma and Rumtse are entirely fed by the glaciers of Kang 406 Yatze region. In 1969, the vegetation area was found to be 17.2 and 2.8 km² in Stok and Kang 407 Yatze region, respectively. Whereas, vegetation cover decreased over time in Stok region as a 408 whole to 14 km² and increased in Kang Yatze region to 3.15 km². However, in few villages of 409 Stok and Kang Yatze region, vegetation increased significantly. These significant changes are 410 discussed in the village-level section (Section 4.2). The change observed over 49 years period 411 was found to be -3.2 and 0.34 km² equivalent to -18.4 (-0.38 %/year) and 12.1 % (0.25 %/year) 412 in Stok and Kang Yatze regions, respectively. 413

Overall, the vegetation cover decreased in the study area from 20.1 km² in 1969 to 17.2 km² in 414 2018. The total change found was -2.8 km² equivalent to -14.1 % (-0.29 %/year) in the entire 415 study area in 49-year period (1969-2018). It should be noted that both the images (Corona and 416 PlanetScope) used in this study are from summer pre-harvesting period to achieve minimum 417 uncertainty. It is also possible that during these particular years (1969 or 2018) the harvest was 418 good due to plenty of water or farming was abandoned due to lack of water, disaster or societal 419 disputes thus raising the level of uncertainty in estimating the actual scenario. We made every 420 possible effort to avoid such issues, and consulted Google Earth temporal image collections 421 and other cloud free PlanetScope imageries of immediate years. 422

Built-up sites of the Stok and Kang Yatze regions increased from 0.61 and 0.15 km² in 1969 to 423 2.27 and 0.3 km² in 2018, with a change in 1.66 and 0.15 km² equivalent to 272 % (5.6 %/year) 424 and 100% (2.1 %/year), respectively. Increase in built-up areas are directly associated with the 425 increase in connectivity of the region. The road network in this region increased significantly 426 with an increase from 81 and 61 km in 1969 to 250 and 105 km in 2018 in Stok and Kang 427 Yatze region, respectively. The rise in road network was found to be 169 km (209 % at 4.3 428 %/year) and 44 km (72 % at 1.5 %/year) in Stok and Kang Yatze region, respectively. A 429 comparatively less change in road network in Kang Yatze region was because all of these 430 431 villages lie directly on the Leh-Manali highway and majority of the road remained same over these years except some link roads within these narrow villages. 432

The villages of Stok region are mostly away from major highways and thus multiple roads 433 linking these villages were made. The mining activity (of gravel for constructions) within the 434 study area are entirely situated in Stok region's alluvial fan, thus giving rise to Built-up area 435 and road network. Villages which were inaccessible for vehicular activity in 1969 are now 436 accessible in 2018. Overall, an increase of 1.81 km² (238 % at 4.9 %/year) and 213 km (150 % 437 at 3.1 %/year) of built-up area and road network was observed in the entire study area, 438 respectively. Figure 5, 6, 7 and 8 presents the built-up area and % change over 49-year period 439 between 1969 and 2018. Local population in the study area has grown more than double its 440 441 size in 50 years at a rate of 2.3 %/year. This has led to an increase in built-up activities in the region. Also, the booming tourism industry with a 600+ fold in tourism footprint (Dame et al., 442 2019; Müller et al., 2020) over the years has initiated over exploitation of natural resources and 443 unplanned urbanisation. The easy money in this industry has caused a shift in local economy 444 from agriculture to tourism thus rendering the once cultivable land more prone to degradation 445 and desertification. 446

Number of lakes increased from 15 lakes (2 in Stok and 13 in Kang Yatze region) in 1969 to 447 36 lakes (5 in Stok and 31 in Kang Yatze region) in 2018. Total lake area of Stok and Kang 448 Yatze regions changed from 0.009 and 0.24 km² in 1969 to 0.02 and 0.52 km² in 2018, 449 respectively. Figure 1, 2, 3 and 5 presents the lake location, area and % change in area in 49 450 years. Overall, change in lake area of the study area was found to be 0.3 km² equivalent to 122 451 % at a rate of 2.5 %/year. Identification of higher altitude lakes formed due to glacial retreat or 452 moraine dammed lake are extremely important in this age when glacial lake outburst floods 453 (GLOFs) are on the rise. GLOFs alone can be devastating especially when unplanned 454 455 urbanisation in the downstream valleys is happening. Flash flood of Kedernath in 2013 and GLOF of Gya valley (Kang Yatze region) in 2014 (Schmidt et al., 2020; Majeed et al., 2021) 456 are some of the few examples of such events. 457

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4.2. Village level change in key environmental parameters

Regional level change of key environmental parameters, already discussed in previous sections, was not sufficient to explain the changes in detail. Therefore, we have also analysed the changes at village-scale. Tuchik, a catchment between Matoo and Martselang catchment, constitute insignificant amount of vegetation cover and no population or settlement area. Though, it is not a village but understanding this catchment is important from a glaciological perspective. This catchment holds only one glacier of Stok region with an area of 0.12 km² (2019) and lost half of its area (0.12 km², 1 %/year) in last 50 years (1969-2019).

Martselang on the other hand also showed similar but comparatively lower change in total area
loss of 38% (0.68 km², 0.8 %/year) followed by Upshi, Matoo and Rumbak catchment with 19
(8.04 km², 0.38 %/year), 18.9 (0.93 km², 0.38 %/year) and 17% (0.38 km², 0.34 %/year) loss
in total glacierised area in 50-year period, respectively. The common element between glaciers

of Tuchik and Martselang catchment is the glacier size i.e., very small (~0.2 km²), therefore 471 the results indicate that glaciers of smaller sizes are more prone to retreat as compared to 472 glaciers with a higher surface area. A small change in equilibrium line altitude or snow line 473 altitude can expose these small glaciers entirely to melting thus losing a relatively higher 474 surface area over time. Whereas glaciers of Stok and Sku-Markha catchment lost comparatively 475 least relative glacierised area of 11.7 % (0.46 km², 0.23 %/year) and 11.4 % (3.48 km², 0.23 476 %/year), respectively. Since, the volume of glacier is related to the glacier area, therefore 477 change in glacier volume is expected to follow the similar trend over 50-year period. The total 478 479 volume change was found to be -38%, -15%, -15%, -21%, -25%, -15% and -12% in Tuchik, Rumbak, Stok, Matoo, Martselang, Upshi and Sku-Markha village catchments, respectively. 480

Built-up change was found to be inversely proportional to the vegetation of the village. Highest 481 increase in built-up area was found in Stok village (688%) followed by Changa (247%), 482 Chuchot (246%), Matoo (225%) and Hemis (205%). The nearly seven-fold rise in built-up area 483 of Stok village is because of the seven-mining activity within the vicinity of village area. Road 484 network in the vicinity of Stok village also increased significantly (380%) and is the highest 485 among all the villages of the study area. Road network of Martselang also increased 486 significantly (368%) followed by Matoo (253%), Changa (141%) and Chuchot (111%). 487 Whereas, comparatively lowest increase in built-up area was found in Sku-Markha, Upshi, and 488 489 Stakna with 62%, 122% and 134 % increase over 50-year period, respectively. Road network in some of these villages are also lowest among the study area with 8% and 28% in Upshi and 490 Stakna. One of the main reasons for such low built-up area and road network is because of the 491 two agricultural farms (one in each Stakna and Upshi) established in recent times (2005), a 492 result of Igu-Phey canal which turned these barren lands into an irrigated farm (GJK, 2011; 493 Nüsser et al 2012). 494

Vegetation of Upshi (48%), Changa (87%) and Rumbak (29.4%) increased whereas a decrease 495 in vegetation was observed in other villages over 40-year period. Increase in vegetation of 496 Upshi and Changa is the outcome of Igu-Phey canal which gave these barren lands an extra 497 source of water for irrigation. Our results are in agreement with GJK, 2011 and Nüsser et al 498 2012. The barren land around Stakna has also been turned into a full-fledged agricultural and 499 animal farm and expected to grow in coming years. Construction of such canals needs to be 500 undertaken to supply water to those regions which has the potential for agricultural activity to 501 increase food production of the region. Increase in built-up area and road network and decrease 502 503 in vegetation is directly linked to the increase in population. Highest population growth was seen in Chuchot, Changa, Stok and Matoo with an increase of 158%, 132%, 112% and 72% in 504 last 50 years (1971-2021). Whereas lowest growth in population was observed in Hemis, 505 Rumbak and Sku-Markha with 33%, 38% and 53%, respectively. Figure 8 presents the change 506 in all the key environmental parameters at a village level between 1969 and 2021. 507

508

509 **5.** Conclusion

510 In this study we analysed key environmental parameters including area and volume of 136 glaciers, built-up and vegetation of 13 villages, area of 36 lakes and demography of 13 major 511 villages around Leh between 1969 and 2021. The area change of glaciers was obtained using 512 freely available Corona and Landsat imageries with additional aid from PlanetScope image and 513 Google Earth. A comparatively higher retreat (~19 at 0.38 %/year) in the glacierised area was 514 found in 50 years as compared to the adjacent region of Shayok and Suru. The retreat was even 515 516 higher (~35-50%) in very small glaciers of area ~0.02 km². Volume of glaciers also decreased over time with a reduction of 0.42 km³ equivalent to \sim 13% (0.8 %/year) of the volume in 16-517 year period. At present, the study area holds 2.5 Gt of water in these glaciers. 518

A decrease (-5 to -58%) in vegetation was found in majority of the villages except Rumbak 519 (29%), Upshi (48%) and Changa (87%) where vegetation significantly increased over 49-vear 520 period. The 43 km long Igu-Phey canal played a major role in such increase in vegetation. 521 Increase in built-up area (688%) and road network (380%) was found highest in Stok village 522 due to seven gravel mining activity in the area whereas, other village witnessed comparatively 523 lower increase in built-up area (63-275%) and road network (8-368%). An inverse relationship 524 was found between % change of vegetation and % change of Built-up area. Increase in 525 population growth percentage was also found highest in Stok (112%), Changa (132%) and 526 527 Chuchot (157%) whereas in other villages the population growth was comparatively lower (32-72%). One of the major reasons for higher increase in built-up areas and road network, and a 528 decrease in vegetation is most probably the tourism industry as tourist footprint in last 44 years 529 was 620 times (527 in 1974 to 327366 tourists in 2018). Another factor for such change can 530 also be attributed to government and defence related activities which is expected to grow in 531 coming years due to redesignation of Ladakh as a Union Territory. Formation of new lakes 532 were also observed in 49 years. Number of lakes increased from 15 to 36 with an increase in 533 area from 0.25 km² to 0.55 km² equivalent to 122% change in lake area between 1969 and 534 2018. 535

While research and policy focus is slowly falling on Leh City and its vicinity, our study proves that environmental changes are happening at a higher rate in rural regions around Leh. A proper assessment and a detailed future planning are very much required in rural regions as well. Systematically documenting these changes using multisource datasets for a larger region can provide more holistic picture of the effects which these mountains are experiencing under the contemporary climate change and demographic shifts.

23 | Page

543 6. Acknowledgement

The authors would like to thank USGS for Corona and Landsat imageries, Planet.com for PlanetScope imagery, NASA for ASTER DEM and The European Centre for Medium-Range Weather Forecasts (ECMWF) for ERA5 reanalysis temperature and precipitation data. We are grateful to Ms. Prerna Joshi for the insights and help in this study. The authors are also thankful to the editor, scientific editor and the two anonymous reviewers for their critical review of the manuscript.

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7. Data availability statement

The data that support the findings of this study are available on request from the lead author.The data belongs to the lead author's PhD program and cannot be made public for now.

553 **8. References**

Andreassen, L. M., Paul, F., Kaab, A., & Hausberg, J. E. (2008). Landsat-derived glacier inventory for Jotunheimen, Norway, and deduced glacier changes since the 1930s. *The Cryosphere*, 16.

Arrowsmith, C., & Inbakaran, R. (2002). Estimating environmental resiliency for the
Grampians National Park, Victoria, Australia: a quantitative approach. Tourism Management,
23(3), 295-309.

Barrett, K., & Bosak, K. (2018). The Role of Place in Adapting to Climate Change: A Case
Study from Ladakh, Western Himalayas. Sustainability, 10(4), 898.
https://doi.org/10.3390/su10040898

Bahr, D. B., Meier, M. F., & Peckham, S. D. (1997). The physical basis of glacier volume-area
scaling. *Journal of Geophysical Research: Solid Earth*, *102*(B9), 20355–20362.
https://doi.org/10.1029/97JB01696

- Bhambri, R, Bolch, T., Kawishwar, P., Dobhal, D. P., Srivastava, D., & Pratap, B. (2013).
 Heterogeneity in glacier response in the upper Shyok valley, northeast Karakoram. *The Cryosphere*, 14.
- 569 Bhambri, Rakesh, Bolch, T., Chaujar, R. K., & Kulshreshtha, S. C. (2011). Glacier changes in
- 570 the Garhwal Himalaya, India, from 1968 to 2006 based on remote sensing. Journal of
- 571 *Glaciology*, 57(203), 543–556. https://doi.org/10.3189/002214311796905604
- Bhardwaj, A., Joshi, P. K., Sam, L., Singh, M. K., Singh, S., & Kumar, R. (2015). Applicability
 of Landsat 8 data for characterizing glacier facies and supraglacial debris. International Journal
 of Applied Earth Observation and Geoinformation, 38, 51-64.
- Bhardwaj, A., Joshi, P. K., Singh, M. K., Sam, L., & Gupta, R. D. (2014). Mapping debriscovered glaciers and identifying factors affecting the accuracy. Cold Regions Science and
 Technology, 106, 161-174.
- Bhardwaj, A., Kumar, R., & Sam, L. (2019). Analysing Geospatial Techniques for Land
 Degradation Studies in Hindu Kush-Himalaya. In Environmental Change in the Himalayan
 Region (pp. 117-135). Springer, Cham.
- Bolch, T., Menounos, B., & Wheate, R. (2010). Landsat-based inventory of glaciers in western
 Canada, 1985–2005. *Remote Sensing of Environment*, 11.
- 583 Bray, J. (1991). Ladakhi history and Indian nationhood. South Asia Research, 11(2), 115-133.
- 584 Buckley, R. C., Pickering, C. M., & Warnken, J. (2000). Environmental Management for
- 585 Alpine Tourism and Resorts in. Tourism and development in mountain regions, 27-45.
- 586 Buntaine, M. T., Mullen, R. B., & Lassoie, J. P. (2007). Human use and conservation planning
- 587 in Alpine areas of Northwestern Yunnan, China. Environment, Development and
- 588 Sustainability, 9(3), 305-324.

- 591 Chen, J. and Ohmura, A.: Estimation of Alpine glacier water resources and their change since
- 592 the 1870s, IAHS Publications Hydrology in Mountainous Regions, I Hydrological
- 593 Measurements; the water cycle, Proceedings of two Lausanne Symposia, August 1990, edited
- 594 by: Lang, H. and Musy, A., IAHS Publ., 193, 127–135, 1990.
- 595 Cuffey, K., & Paterson, W. S. B. (2010). *The physics of glaciers* (4th ed). Butterworth-596 Heinemann/Elsevier.
- Dame, J., Schmidt, S., Müller, J., & Nüsser, M. (2019). Urbanisation and socio-ecological
 challenges in high mountain towns: insights from Leh (Ladakh), India. Landscape and urban
 planning, 189, 189-199.
- Draper, D. (2000). Toward sustainable mountain communities: Balancing tourism
 development and environmental protection in Banff and Banff National Park, Canada.
 AMBIO: A Journal of the Human Environment, 29(7), 408-415.
- Frey, H., Machguth, H., Huss, M., Huggel, C., Bajracharya, S., Bolch, T., Kulkarni, A.,
 Linsbauer, A., Salzmann, N., & Stoffel, M. (2014). Estimating the volume of glaciers in the
 Himalayan–Karakoram region using different methods. *The Cryosphere*, 8(6),
 2313–2333. https://doi.org/10.5194/tc-8-2313-2014
- 607 GJK (Government of Jammu and Kashmir) 2011. Evaluation Report On Igophey Canal, Leh
- 608 Ladakh (Joint Vernture of Irrigation Division Igophey and CAD Leh) (1979 2011).
- 609 Directorate of Economics and Statistics, J&K Planning and Development Department.

Garg, P. K., Shukla, A., & Jasrotia, A. S. (2019). On the strongly imbalanced state of glaciers
in the Sikkim, eastern Himalaya, India. *Science of The Total Environment*, 691, 16–35.
https://doi.org/10.1016/j.scitotenv.2019.07.086

Geneletti, D., & Dawa, D. (2009). Environmental impact assessment of mountain tourism in
developing regions: A study in Ladakh, Indian Himalaya. Environmental impact assessment
review, 29(4), 229-242.

Haeberli W, Hölzle M (1995) Application of inventory data for estimating characteristics of
and regional climate-change effects on mountain glaciers: a pilot study with the European Alps.
Annals of Glaciology 21: 206-212.

Hoelzle, M., Chinn, T., Stumm, D., Paul, F., Zemp, M., & Haeberli, W. (2007). The application
of glacier inventory data for estimating past climate change effects on mountain glaciers: A
comparison between the European Alps and the Southern Alps of New Zealand. *Global and Planetary Change*, 56(1–2), 69–82. https://doi.org/10.1016/j.gloplacha.2006.07.001

Ines, A. V. M., & Hansen, J. W. (2006). Bias correction of daily GCM rainfall for crop
simulation studies. *Agricultural and Forest Meteorology*, 10.

Kanda, N., Negi, H. S., Rishi, M. S., & Kumar, A. (2020). Performance of various gridded
temperature and precipitation datasets over Northwest Himalayan Region. *Environmental Research Communications*, 2(8), 085002. https://doi.org/10.1088/2515-7620/ab9991

Khan, A., Naz, B. S. & Bowling, L. C. Separating snow, clean and debris covered ice in the Upper
Indus Basin, Hindukush-Karakoram-Himalayas, using Landsat images between 1998 and 2002. *Journal of Hydrology* 521, 46-64 (2015).

Linsbauer, A., Paul, F., & Haeberli, W. (2012). Modeling glacier thickness distribution and
bed topography over entire mountain ranges with GlabTop: Application of a fast and robust

- approach: REGIONAL-SCALE MODELING OF GLACIER BEDS. Journal of Geophysical
- 634 *Research: Earth Surface*, *117*(F3), n/a-n/a. https://doi.org/10.1029/2011JF002313
- Linsbauer, A., Paul, F., Hoelzle, M., Frey, H., & Haeberli, W. (2009). *The Swiss Alps Without Glaciers A GIS-based Modelling Approach for Reconstruction of Glacier Beds*. 6.
- Majeed, U., Rashid, I., Sattar, A., Allen, S., Stoffel, M., Nüsser, M., Schmidt S. (2021)
 Recession of Gya Glacier and the 2014 glacial lake outburst flood in the Trans-Himalayan
 region of Ladakh, India. Science of the Total Environment, 756:144008. doi:
 10.1016/j.scitotenv.2020.144008
- Messerli, B., & Ives, J. D. (1997). Mountains of the world: a global priority. Parthenon, NewYork.
- Michaud, J. (1991). A social anthropology of tourism in Ladakh, India. Annals of Tourism
 Research, 18(4), 605-621.
- Moss, L. A., & Godde, P. M. (2000). Strategy for future mountain tourism. Tourism and
 development in mountain regions., 323-338.
- 647 Müller, J.; Dame, J.; Nüsser, M. Urban Mountain Waterscapes: The Transformation of Hydro-
- Social Relations in the Trans-Himalayan Town Leh, Ladakh, India. *Water* 2020, *12*, 1698.
 https://doi.org/10.3390/w12061698
- 650 Muhammad, S., Gul, C., Muneer, J. & Waqar, M. M. Efficiency of classification techniques
- 651 for sachen and rupal glaciers variation during 1972–2010. International Conference on
- Aerospace Science & Engineering (ICASE), IEEE International, 1-3 (2013).
- Muhammad, S., Tian, L., & Khan, A. (2019). Early twenty-first-century glacier mass losses in
- the Indus Basin constrained by density assumptions. Journal of Hydrology, 574, 467-475.

- Nagai, H., Fujita, K., Sakai, A., Nuimura, T. & Tadono, T. Comparison of multiple glacier
 inventories with a new inventory derived from high-resolution ALOS imagery in the Bhutan
 Himalaya. The Cryosphere 10, 65-85, doi:10.5194/tc-10-65-2016 (2016).
- Negi, H. S., Kumar, A., Kanda, N., Thakur, N. K., & Singh, K. K. (2021). Status of glaciers
- and climate change of East Karakoram in early twenty-first century. *Science of The Total*
- 660 Environment, 753, 141914. https://doi.org/10.1016/j.scitotenv.2020.141914
- Nepal, S. K., & Chipeniuk, R. (2005). Mountain tourism: Toward a conceptual framework.
 Tourism Geographies, 7(3), 313-333.
- Nüsser, M., Dame, J., Kraus, B., Baghel, R., & Schmidt, S. (2019). Socio-hydrology of
 "artificial glaciers" in Ladakh, India: Assessing adaptive strategies in a changing cryosphere.
 Regional Environmental Change, 19(5), 1327–1337. https://doi.org/10.1007/s10113-0181372-0
- 667 Nüsser, M., Schmidt, S., & Dame, J. (2012). Irrigation and Development in the Upper Indus
- 668 Basin: Characteristics and Recent Changes of a Socio-hydrological System in Central Ladakh,
- 669 India. Mountain Research and Development, 32(1), 51-61. https://doi.org/10.1659/MRD-
- 670 JOURNAL-D-11-00091.1
- Nuth, C., & Kääb, A. (2011). Co-registration and bias corrections of satellite elevation data
 sets for quantifying glacier thickness change. *The Cryosphere*, 5(1), 271–290.
 https://doi.org/10.5194/tc-5-271-2011
- Paul, F., & Kääb, A. (2005). Perspectives on the production of a glacier inventory from
 multispectral satellite data in Arctic Canada: Cumberland Peninsula, Baffin Island. *Annals of Glaciology*, 42, 59–66. https://doi.org/10.3189/172756405781813087

- 677 Price, M. F. (1992). Patterns of the development of tourism in mountain environments.678 GeoJournal, 27(1), 87-96.
- 679 Ramsankaran, R., Pandit, A., & Azam, M. F. (2018). Spatially distributed ice-thickness

modelling for Chhota Shigri Glacier in western Himalayas, India. International Journal of

- 681 *Remote Sensing*, *39*(10), 3320–3343. https://doi.org/10.1080/01431161.2018.1441563
- 682 Salzmann, N., Huggel, C., Rohrer, M., Silverio, W., Mark, B. G., Burns, P., & Portocarrero, C.
- 683 (2013). Glacier changes and climate trends derived from multiple sources in the data scarce
- 684 Cordillera Vilcanota region, southern Peruvian Andes. The Cryosphere, 7(1), 103–118.
- 685 https://doi.org/10.5194/tc-7-103-2013
- Sam, L., Kumar, R., & Bhardwaj, A. (2019). Climate and Remotely Sensed Markers of Glacier
 Changes in the Himalaya. In Environmental Change in the Himalayan Region (pp. 65-88).
 Springer, Cham.
- Schmidt, S., & Nüsser, M. (2017). Changes of High Altitude Glaciers in the Trans-Himalaya
 of Ladakh over the Past Five Decades (1969–2016). *Geosciences*, 7(2), 27.
 https://doi.org/10.3390/geosciences7020027
- 692 Schmidt, S., Nüsser, M., Baghel, R., & Dame, J. (2020). Cryosphere hazards in Ladakh: The
- 693 2014 Gya glacial lake outburst flood and its implications for risk assessment. *Natural Hazards*,
- 694 104(3), 2071–2095. https://doi.org/10.1007/s11069-020-04262-8
- Sen, P. K.: Estimates of the regression coefficient based on Kendall's Tau, Am. Stat. J., 63,
 1379–1389, https://doi.org/10.2307/2285891, 1968.
- 697 Shrestha, M., Acharya, S. C., & Shrestha, P. K. (2017). Bias correction of climate models for
- 698 hydrological modelling are simple methods still useful? *Meteorol. Appl.*, 9.

699	Shukla, A	A., & Garg	, P. K.	(2019).	Evolution of a c	lebris-covered glacies	r in the v	western
700	Himalaya	during the	e last fo	our decad	les (1971–2016):	A multiparametric a	issessmen	t using
701	remote	sensing	and	field	observations.	Geomorphology,	341,	1–14.
702	https://do	i.org/10.101	16/j.geoi	morph.20	19.05.009			

703 Shukla, A., Garg, S., Mehta, M., Kumar, V., & Shukla, U. K. (2020). Temporal inventory of

glaciers in the Suru sub-basin, western Himalaya: Impacts of regional climate variability. Earth

- 705 System Science Data, 12(2), 1245–1265. https://doi.org/10.5194/essd-12-1245-2020
- Smith, S. (2012). Intimate geopolitics: Religion, marriage, and reproductive bodies in Leh,
- Ladakh. Annals of the Association of American Geographers, 102(6), 1511-1528.
- Soheb, M., Ramanathan, A., Angchuk, T., Mandal, A., Kumar, N., & Lotus, S. (2020). Mass-708 709 balance observation, reconstruction and sensitivity of Stok glacier, Ladakh region, India, 1978 2019. Glaciology, 710 between and Journal 66(258), 627-642. of 711 https://doi.org/10.1017/jog.2020.34
- Stevens, S. (2003). Tourism and deforestation in the Mt Everest region of Nepal. Geographical
 Journal, 169(3), 255-277.
- Teutschbein, C., & Seibert, J. (2013). Is bias correction of regional climate model (RCM)
 simulations possible for non-stationary conditions? *Hydrology and Earth System Sciences*, *17*(12), 5061–5077. https://doi.org/10.5194/hess-17-5061-2013
- Tian, L. et al. Two glaciers collapse in western Tibet. Journal of Glaciology 63, 194-197,
 doi:10.1017/jog.2016.122 (2017).
- Warikoo, K. (2020). Ladakh: India's Gateway to Central Asia. Strategic Analysis, 44(3), 177192.

Table 1: Presents the data used in the present study.

	Data	Scene ID	Years	Spatial res.	Use
1	Landsat TM Landsat TM Landsat TM Landsat OLI	LT05_L1TP_147037_19910828_20170126_01_T1 LT05_L1TP_147037_19980916_20161222_01_T1 LT05_L1TP_147037_20100917_20161013_01_T1 LC08_L1TP_147037_20190910_20190917_01_T1	1991, 1998, 2010, 2019	30m	Glacier mapping
2	Corona	DS1107-1104DA015 (30 July) DS1107-1104DA016 (30 July) DS1107-1104DA017 (30 July) DS1107-1104DA018 (30 July) DS1107-1104DA019 (30 July) DS1107-1104DA020 (30 July)	1969	3m	Glacier, vegetation, built-up and water body mapping
3	ASTER DEM	AST14DMO.003:2025691837 (08 September) AST14DMO.003:2025691835 (08 September) AST14DMO.003:2404575282 (20 September) AST14DMO.003:2404575284 (20 September)	2004 2004 2020 2020	30m	Glacier thickness estimation, catchment delineation.
4	PlanetScope	20180829_045714_0f2a, 20180829_045715_0f2a, 20180829_045716_0f2a, 20180829_045717_0f2a, 20180829_045718_0f2a, 20180829_045719_0f2a, 20180829_045722_0f2a, 20180829_045723_0f2a, 20180829_045722_0f2a, 20180829_045725_0f2a, 20180829_045726_0f2a, 20180829_045725_0f2a, 20180829_045726_0f2a, 20180829_050228_103b, 20180829_050229_103b, 20180829_050232_103b, 20180829_050231_103b, 20180829_050232_103b, 20180829_050233_103b, 20180829_050308_103d, 20180829_050309_103d, 20180829_050310_103d, 20180829_050311_103d, 20180829_050312_103d, 20180829_050315_103d, 20180829_050314_103d, 20180829_050315_103d, 20180829_050316_103d, 20180829_050317_103d, 20180829_050318_103d, 20180829_050319_103d, 20180829_050318_103d, 20180829_050319_103d, 20180829_050318_103d, 20180829_050319_103d, 20180829_050318_103d, 20180829_050319_103d, 20180829_050318_103d, 20180829_050319_103d, 20180829_050318_103d, 20180829_050319_103d, 20180829_050318_103d, 20180829_050319_103d, 20180829_050318_103d, 20180829_050319_103d, 20180829_050422_0f17, 20180829_050424_0f17, 20180829_050422_0f17, 20180829_050424_0f17, 20180829_050422_0f17, 20180829_050424_0f17, 20180829_050422_0f17, 20180829_050424_0f17, 20180829_050425_0f17, 20180829_050428_0f17, 20180829_050425_0f17, 20180829_050428_0f17, 20180829_050429_0f17, 20180829_050428_0f17, 20180829_050429_0f17, 20180829_050428_0f17, 20180829_050429_0f17, 20180829_050532_0f28, 20180829_050429_0f17, 20180829_050532_0f28, 20180829_050533_0f28, 20180829_050534_0f28, 20180829_050535_0f28	2018	3m	Glacier, vegetation, built-up and water body mapping
5	ERA5 data	ERA5 hourly data on single levels from 1979 to 2019. DOI: 10.24381/cds.adbb2d47	1979- 2019	0.25°	Climate analysis
6	Census data	Census report of 1971, 1981, 2001 and 2011.	1971, 1981, 2001, 2011	10 year	Demographic change

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		ANNUAL		JJAS	JAS				FER	-			
ERA5 C points	Grid	Mann-Kendall Significance (p)	Sens Slope (β)	Long-term Mean (°C or mm)	Change (°C or mm)	Mann-Kendall Significance (p)	Sens Slope (β)	Long-term Mean (°C or mm)	Change °C or mm	Mann-Kendall Significance (p)	Sens Slope (β)	Long-term Mean (°C or mm)	Change °C or mm
Leh Tempera	ture	0.01	0.04	7.65	0.21	0.01	0.05	18.67	0.11	0.05	0.03	-0.45	-2.73
Leh Precipita	tion	0.4	-0.33	58.12	-0.23	0.6	-0.1	27.65	-0.15	0.9	0.001	19.27	0.00

Table 2: Results of Mann -Kendall and Sen's slope estimator for annual, JJAS and winter

temperature and precipitation.

D .	1969 1991		1998	2010	2019	
Region	km ²	km^2	km ²	km ²	km ²	
Stok	±0.1 (0.8%)	±0.6 (6.1%)	±0.6 (6.1%)	±0.6 (6.6%)	±0.6 (7.1%)	
Kang Yatze	±0.5 (0.7%)	±3.4 (4.9%)	±3.3 (5.0%)	±3.2 (5.4%)	±3.1 (5.6%)	
All	±0.6 (0.5%)	±4 (3.9%)	±3.9 (4.0%)	±3.8 (4.1%)	±3.7 (4.2%)	

Table 3: Uncertainty in glacierised area for different years through buffer based assessment.

- Table 4: Estimated uncertainty in mean thickness (m) and total ice volumes (km³) in Glabtop2,
- slope dependent and area scaling methods.

Uncertainties in glabtop2 and slope dependent methods									
0		20	04		2020				
Ŋ	Thic	kness	Volume		Thicknes	S S	Volume		
RI	n	n	km ³		m		km ³		
Stok	±4.6 (±	23.7%)	±0.002 (±23.8%)	±3.8 (±22.	6%) ±	-0.002 (±22	.6%)	
Kang Yatze	±6.4 (±	25.5%)	±0.009 (±25.7%)		±4.9 (±23	%) ±	=0.006 (±23	.1%)	
All	±6 (±2	25.1%)	±0.007 (±25.3%)	±4.63 (±23	3%) ±	0.005 (±22	.9%)	
	I		Uncertain	ties in area s	caling method				
		20	04		2020				
NO NO	Glacier a	area +5%	Glacier area -5% Glacier area			ea +5% Glacier area -5%			
REGI	Chen and Ohmura (1990)Bahr et al., (1997)		Chen and Ohmura (1990)	Bahr et al., (1997)	Chen and Ohmura (1990)	Bahr et al., (1997)	Chen and Ohmura (1990)	Bahr et al., (1997)	
	km ³	km ³	km ³	<i>km³</i>	km ³	km ³	km ³	km ³	
	(Differenc	(Differenc	(Differenc	(Differenc	(Difference)	(Difference	(Differenc	(Differ	
Stok	<i>e)</i>	<i>e)</i>	<i>e)</i>	<i>e)</i>	0.27(+8.0%))	e_{j}	ence)	
SIOK	$\begin{bmatrix} 0.51 & 0.55 & 0.27 & 0.51 \\ (+6.9\%) & (+6.1\%) & (-6.9\%) & (-6.1\%) \end{bmatrix}$		(-6.1%)	0.27(+0.070)	(+7.1%)	(-8.0%)	(-7.1%)		
Kang	3.08	3.55	2.69	3.09	2.85 (+6.7%)	3.29	2.5	2.87	
Yatze	(+6.9%)	(+6.9%)	(-6.6%)	(-6.9%)	· · · ·	(+7.2%)	(-6.4%)	(-6.5%)	
All	3.39	3.9	2.96	3.4	3.12 (+6.8%)	3.59	2.73	3.13	
	(+6.9%)	(+6.8%)	(-6.6%)	(-6.8%)		(+7.2%)	(-6.5%)	(-6.6%)	

Uncertainties in the lake area									
	Anal	yst A	Analyst B		Ana	lyst C	±Uncertainty (%)		
REGION	1969 (km2)	2018 (km2)	1969 (km2)	2018 (km2)	1969 (km2)	2018 (km2)	1969	2018	
Stok Region	0.009	0.022	0.009	0.021	0.008	0.023	2.35	2.61	
Kang Yatze Region	0.237	0.524	0.231	0.511	0.239	0.523	1.33	1.17	
All	0.246	0.546	0.240	0.532	0.247	0.546	1.21	1.22	
			Uncertain	ties in the b	uild-up area	<u>l</u>			
SAMPLE VILLAGE	1969 (m2)	2018 (m2)	1969 (m2)	2018 (m2)	1969 (m2)	2018 (m2)	1969	2018	
Chuchot Village	0.280	0.969	0.264	0.939	0.273	0.989	2.28	2.12	
Stok Village	0.062	0.492	0.061	0.517	0.064	0.501	1.91	2.08	
All	0.342	1.461	0.325	1.456	0.337	1.490	2.06	1.04	
	Accutacy assessment for the vegetation area								
DECION		1	969		2018				
REGION	Overall A	ccuracy	Карра С	officient	Overall .	Overall Accuracy		officient	
Stok	79%			0.72	92%		0.9		
Kang Yatze	Kang Yatze 74%			0.6	89%		0.86		
All 77%			0.66		91%		0.88		

Table 5: Uncertainty in lake and build up area, and accuracy assessment for vegetation area.



Figure 1: Location map of the study area presenting all the mapped key environmental
parameters. Advanced Spaceborne Thermal Emission and Reflection Radiometer Digital
Elevation Model (ASTER DEM) has been used in the background to provide the topographic
details. The inset map highlights the contextual location of study area in North India.



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- **Figure 2**: Presents the glacier outlines of 1969 and 2019, glacier thickness of 2020 and location
- 789 of lakes in Stok region



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Figure 3: Presents the glacier outlines of 1969 and 2019, glacier thickness of 2020 and location

793 of lakes in Kang Yatze region



Figure 4: ERA5 bias corrected temperature and precipitation of Leh grid from 1979 to 2019.
a and b are the correlation between observed and ERA5 reanalysis temperature and
precipitation. c and d are annual and monthly temperature and precipitation.



Figure 5: Total change of key environmental parameters over the years in the study area, Stok





Figure 6: Presents Vegetation, Built-up and Waterbodies map over the years in Stok region.





Figure 7: Presents Vegetation, Built-up and Waterbodies map over the years in Kang Yatze

814 region

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Figure 8: Village level change of key environmental parameters over the years in the studyarea. The section with pink background represents the Indus-fed villages

825	Figure 9: Location map of the study area presenting all the mapped key environmental
826	parameters. Advanced Spaceborne Thermal Emission and Reflection Radiometer Digital
827	Elevation Model (ASTER DEM) has been used in the background to provide the topographic
828	details. The inset map highlights the contextual location of study area in North India.
829	Figure 10: Presents the glacier outlines of 1969 and 2019, glacier thickness of 2020 and
830	location of lakes in Stok region
831	Figure 11: Presents the glacier outlines of 1969 and 2019, glacier thickness of 2020 and
832	location of lakes in Kang Yatze region
833	Figure 12: ERA5 bias corrected temperature and precipitation of Leh grid from 1979 to 2019.
834	a and b are the correlation between observed and ERA5 reanalysis temperature and
835	precipitation. c and d are annual and monthly temperature and precipitation.
836	Figure 13: Total change of key environmental parameters over the years in the study area,
837	Stok and Kang Yatze regions
838	Figure 14: Presents Vegetation, Built-up and Waterbodies map over the years in Stok region.
839	Figure 15: Presents Vegetation, Built-up and Waterbodies map over the years in Kang Yatze
840	region
841	Figure 16: Village level change of key environmental parameters over the years in the study
842	area. The section with pink background represents the Indus-fed villages